

INTERPRETATION OF REMOTE SENSING DATA  
FROM EASTERN AREAS OF THE BALTIC SEA

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A method for the determination of concentrations of optically active substances (phytopigments, suspended and yellow matter) in waters with high concentrations and unequal indices of optically active substances is suggested. The method has been worked out relying on the Baltic Sea phenomenon but it can also be applied to other water bodies (seas as well as internal ones) of similar parameters.

Possible ways of applying remote sensing to the admixture concentration determination in the eastern Baltic waters are considered. It is known that the waters there are highly trophic and strongly subjected to human impact. In most papers devoted to the interpretation of spectral aerocosmic information on the sea surface, empiric relations between one or two characteristics of upward radiation spectrum and the chlorophyll "a" concentration were reached (Clark, 1981; Smith and Baker 1982; Shifrin, 1983, Shturm, 1984; Sugihara, Kishino and Okami, 1985). In the present paper an attempt of a more profound use of the upward radiation spectrum has been made. Let us consider the coefficient of the sea surface spectral brightness equal to

$$R = \frac{L_{s\lambda} - L_{z\lambda}}{L_{o\lambda}} \quad (\text{Lokk and Pelevin, 1978; Pelevin, 1978}), \text{ where}$$

$L_{s\lambda}$  is the brightness of the upward radiation above the sea surface,  $L_{z\lambda}$  the sky brightness in the zenith,  $L_{o\lambda}$  the brightness of the white diffuser, placed horizontally above the sea surface,  $f = 0.02$  the value of the Frenel reflection coefficient for the smooth water surface and normal light beam inclination,  $\lambda$  the light wavelength. To determine the spectra of  $\rho_{\lambda}$ ,  $L_{s\lambda}$ ,  $L_{z\lambda}$  and  $L_{o\lambda}$  were directly measured. The experimental values of  $\rho_{\lambda}$  were obtained in the Baltic in 1986 and 1987 using the spectroradiometer with a set of interferential lightfilters:

$$\begin{aligned} \lambda_1 &= 381 \text{ nm}; & \lambda_2 &= 438 \text{ nm}; & \lambda_3 &= 500 \text{ nm}; \\ \lambda_4 &= 555 \text{ nm}; & \lambda_5 &= 596 \text{ nm}; & \lambda_6 &= 670 \text{ nm}. \end{aligned}$$

For calculations the formula of spectral brightness coefficient most directly connecting the parameter with the hydrooptical characteristic of the water, was chosen (Lokk and Pelevin, 1978):

$$\rho_{\lambda} = \kappa \frac{b_{\beta\lambda}}{a_{\omega\lambda} + a_{y\lambda} + a_{p\lambda} + a_0 + b_{\beta\lambda}}, \quad (1)$$

where  $b_{\beta\lambda}$  is the coefficient of light backscattering;  $a_{\omega\lambda}$ ,  $a_{p\lambda}$ ,  $a_{y\lambda}$ ,  $a_0$  the coefficients of absorption in pure water, plankton pigments, dissolved organic "yellow" matter, suspended matter (excluding pigments, phytoplankton),  $\kappa = 0.11$  (Pelevin, 1978). It is believed (Pelevina, 1980) that the coefficient of absorption in the dissolved organic ("yellow") matter changes with the spectrum as  $a_{y\lambda} = a_{y\lambda_0} \cdot \exp K(\lambda_0 - \lambda)$ , where  $K = 0.0148 \div 0.0152 \text{ m}^{-1}$ . In calculations  $K = 0.015 \text{ m}^{-1}$  and  $\lambda = 500 \text{ nm}$ . It is presumed that the construction of absorption in phytoplankton pigments is proportional to the chlorophyll "a" concentration  $C \text{ mg} \cdot \text{m}^{-3}$ :

$$a_{p\lambda} = a_{p\lambda}^* \cdot C \text{ m}^{-1}.$$

After the analysis of various publications (Jerlov, 1970; Ochakovskij, et al., 1974; Popov, Fedorov and Orlov, 1979; Optics of Ocean, 1983) the following coefficient values were adopted:

$$a_{p381}^* = 0.036 \text{ m}^2 \cdot \text{mg}^{-1}; \quad a_{p438}^* = 0.059 \text{ m}^2 \cdot \text{mg}^{-1};$$

$$a_{p500}^* = 0.042 \text{ m}^2 \cdot \text{mg}^{-1} \text{ for specific absorption in pigments}$$

$$\text{and } a_{\omega} = 0.014 \text{ m}^{-1}; \quad a_{\omega438} = 0.011 \text{ m}^{-1}; \quad a_{\omega500} = 0.021 \text{ m}^{-1};$$

$$a_{\omega555} = 0.053 \text{ m}^{-1}; \quad a_{\omega596} = 0.15 \text{ m}^{-1} \text{ in pure water (on natural}$$

basis). The spectral dependence  $b_{\beta\lambda}$  in condition of the additive contribution of backscattering by large  $b_{\beta\lambda}$  and small  $b_{\beta\lambda}$  fractions of suspended matter is selected as follows (Optics of Ocean, 1983):

$$b_{\beta\lambda} = b_{\beta\omega500} \cdot \left(\frac{500}{\lambda}\right)^{4.3} + b_{\beta\lambda500} \cdot \left(\frac{500}{\lambda}\right)^{0.3} + b_{\beta5500} \cdot \left(\frac{500}{\lambda}\right)^{1.7}.$$

The coefficient of light backscattering in pure water is  $b_{\beta\omega500} = 0.001 \text{ m}^{-1}$  (Popov, Fedorov and Orlov, 1979). It was suggested that within the interval of  $500 \text{ nm} \leq \lambda \leq 600 \text{ nm}$  the special absorption in pigments  $a_{p\lambda}^*$  changes according to the law  $\exp[0.015(\lambda_0 - \lambda)]$ . The light absorption in suspended matter  $a_0$  (excluding the absorption in pigments) for calculations is considered non-selective. This way, the light absorption in pigments and dissolved yellow matter  $a_{\lambda}$  changes on the spectrum

$$a_{\lambda} = a_{y\lambda} + a_{p\lambda} = a_{500} \cdot \exp[0.015 \cdot (500 - \lambda)].$$

At wavelengths  $500 \text{ nm} \leq \lambda \leq 600 \text{ nm}$  equation (1) is expressed as follows:

$$\rho_{\lambda} = \kappa \frac{b_{\beta\lambda}}{a_{\omega\lambda} + a_{500} \cdot \exp[0.015(500 - \lambda)] + a_0 + b_{\beta\lambda}} \quad (2)$$

For  $\lambda < 500 \text{ nm}$ , where the spectral way of the curves  $a_{y\lambda}$  and  $a_{p\lambda}$  differs considerably, equation (1) has the form

$$\rho_{\lambda} = \kappa \frac{b_{\beta\lambda}}{a_{\omega\lambda} + (1 - \xi) \cdot a_{500} \cdot \exp[0.015(500 - \lambda)] + \xi \cdot a_{500} \cdot N_{\lambda} + b_{\beta\lambda}}, \quad (3)$$

where  $\xi = \frac{a_{p500}}{a_{500}}$ ;  $N_{\lambda} = \frac{a_{p\lambda}^*}{a_{p500}^*}$

The reverse task - the determination of the unknown parameters  $b_{\lambda 500}$ ,  $b_{\lambda 550}$ ,  $a_{y500}$ ,  $a_{p500}$ ,  $a_0$  for measuring  $\rho_{\lambda i}$ , was solved by a computer, relying on the principle of minimizing the discrepancies of the measured values of  $\rho_{\lambda i}$  and of the calculated values of  $\rho'_{\lambda i}$  obtained from (2) - (3). The procedure of calculation is as follows. Firstly the calculation values of  $b_{\lambda 500}$ ,  $b_{\lambda 550}$ ,  $a_{500}$ ,  $a_0$  for  $\sum_{i=3}^5 |\rho_{\lambda i} - \rho'_{\lambda i}| = 0$  are found according to linear equations system (2). Simultaneously according to system (3) by way of the modification of  $a_{500}$  and  $\xi$ , the values of  $a_{p500}$  and  $a_{y500}$  for which the maximum values of the difference modulus  $\sum_{i=4}^2 |\rho_{\lambda i} - \rho'_{\lambda i}|$  are the smallest, are determined. The following conditions are satisfied:

$$b_{\lambda 500}, b_{\lambda 550}, a_{500}, a_0 \geq 0; \quad 0 \leq \xi \leq 1.$$

The results of calculations and measurements for stations showed in Fig. 1, are summarized in Table .

As one can see, the calculated values of the chlorophyll concentration are in good agreement with the concentrations measured in samples. The precision achieved meets totally the economic needs as well as the requirements for the studies of the large-scale variability of water mass and substance distribution in coastal areas. An analogous precision was reached for other optically active substances (suspended and yellow matter) in the water. The algorithm like that may be applied to investigations from low-flying carriers, where the influence of the atmosphere is inessential.

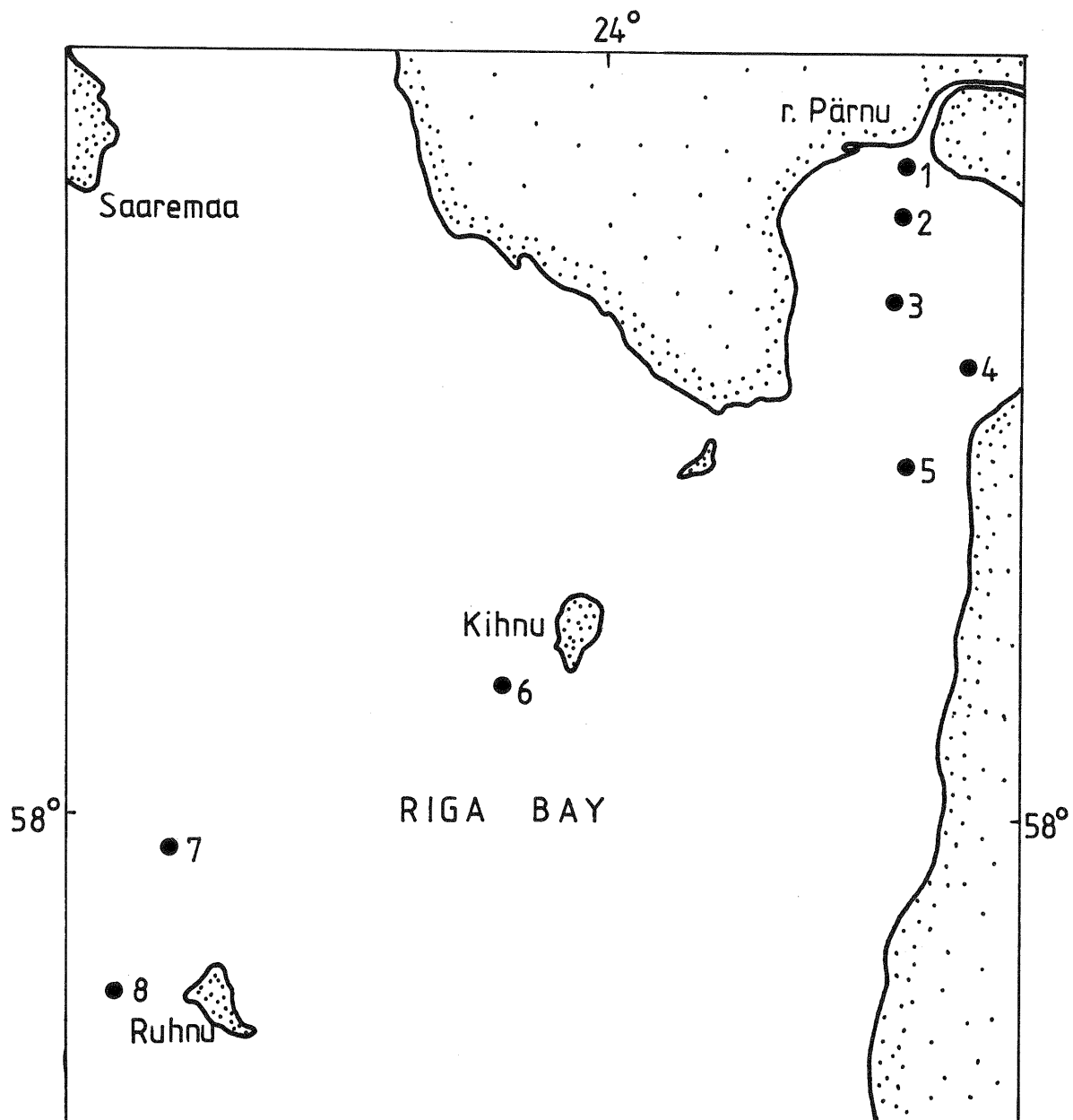


Fig.1. The area of sampling and remote measurements

The measured and calculated values of optical parameters  
and chlorophyll concentration

Table

Station		$\lambda$ (nm)					$m^{-1}$					$C$ ( $mg \cdot m^{-3}$ )	
		383	438	500	555	596	$a_{500}$	$a_0$	$a_{45500}$	$b_{ll500}$	$b_{s500}$	calc.	meas.
1	measured	0.007	0.016	0.026	0.040	0.045	1.10	0.11	0.80	0.38	0.01	7.3	9.0
	calculated	0.008	0.015	0.026	0.040	0.045							
2	measured	0.013	0.022	0.036	0.054	0.062	1.56	0.06	1.00	0.76	0.04	13.2	9.2
	calculated	0.013	0.022	0.036	0.054	0.062							
3	measured	0.015	0.020	0.038	0.056	0.061	0.91	0.01	0.66	0.49	0.01	6.1	7.3
	calculated	0.013	0.022	0.038	0.056	0.061							
4	measured	0.008	0.012	0.020	0.033	0.037	0.89	0.02	0.44	0.19	0.02	10.8	8.5
	calculated	0.008	0.012	0.020	0.033	0.037							
5	measured	0.006	0.007	0.014	0.023	0.025	0.75	0.06	0.42	0.12	0.00	7.9	8.7
	calculated	0.005	0.008	0.014	0.023	0.025							
6	measured	0.003	0.005	0.007	0.010	0.008	0.35	0.01	0.20	0.00	0.03	3.5	2.8
	calculated	0.003	0.005	0.007	0.010	0.008							
7	measured	0.004	0.003	0.007	0.012	0.012	0.52	0.02	0.23	0.04	0.00	6.9	2.7
	calculated	0.003	0.004	0.007	0.012	0.012							
8	measured	0.003	0.003	0.005	0.007	0.006	0.30	0.08	0.12	0.02	0.00	4.2	3.2
	calculated	0.003	0.003	0.005	0.007	0.006							

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